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# Multiscale predictability of Cutaneous Leishmaniasis in Morocco and Tunisia through the AMO-NAO coupling and its modulation of regional rainfall

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## **Abstract**

The development of effective Early Warning Systems (EWS) for climate-driven zoonotic diseases has been hindered by a lack of predictors with adequate lead time for effective interventions. Atmosphere-Ocean coupled phenomena present predictability beyond the atmospheric deterministic limits and therefore are potentially useful climate drivers to be integrated in mathematical models. While the El Niño-Southern Oscillation (ENSO) has been used to forecast disease dynamics in equatorial and tropical regions, there is a lack of similar applications for temperate areas, likely because of the perceived unpredictability of atmospheric systems such as the North Atlantic Oscillation (NAO). This study challenges this notion by establishing a connection between the NAO and its oceanic counterpart, the Atlantic Multidecadal Oscillation (AMO), revealing common low-frequency components that strongly modulate Cutaneous Leishmaniasis (CL) in Northern Africa. We demonstrate not only short-term couplings, such as the known NAO's impact on seasonal rainfall, which subsequently affects CL incidence, but we also uncover a significant lagged effect of approximately three years on rainfall and four years on CL incidence. Our findings reveal a unified, multiscale mechanism that influences CL epidemiology across different time scales, underscoring the predictive skill for short and long term time frames, which should be integrated in CL forecasting models.

## **Introduction**

### Atmosphere-Ocean coupled phenomena and Early Warning Systems

Climate exerts a significant influence on many infectious diseases. In particular, most vector-borne diseases are influenced by prevailing weather conditions, with factors such as extrinsic incubation times, mortality and biting rates, among others, all showing sensitivity to climate variations (Caminade et al 2018). Further, the epidemiology of zoonotic diseases, known for their intricate and often delayed responses to climate, can be influenced through a multitude of complex pathways, mediated by climate-induced alterations in the landscape and vegetation, thereby impacting the population dynamics of reservoir and vector species.

Climate forcing operates across various temporal and spatial scales. Historically, infectious disease ecologists predominantly focused on examining the responses of their phenomenon of interest to local weather patterns (Stenseth et al, 2003). If strong associations emerged, this created the opportunity to make predictions days, or at a maximum weeks, ahead, depending on how slow the disease system responded to this weather forcing. As it is well known, the predictability of the atmospheric system is theoretically constrained by very short deterministic limits, making the extension of such forecasts to larger lags impossible (Tribbia et al, 1993). However, some large-scale climate processes are coupled to other more slowly-evolving systems having greater inertia (i.e. predictability), mainly in the ocean

57 compartment. An illustrative example is the ocean-atmosphere coupled phenomenon known as the El  
58 Niño-Southern Oscillation (ENSO), which under certain conditions creates windows of predictability  
59 months (and years) in advance (Petrova et al, 2021). In fact, ENSO has already been demonstrated to be  
60 able to impact the interannual dynamics of different tropical infectious diseases, allowing for the much-  
61 needed long-term predictability (Pascual et al, 2000). Under these circumstances, extended  
62 predictability beyond the deterministic limits can be observed (Chen et al, 2020).

63  
64 Even when theoretically possible and of potential interest to many regions and diseases, there is a  
65 notable scarcity of operational early warning systems (EWS) designed to function on the necessary long  
66 time-scales required and providing sufficient time for effective intervention strategies. While some  
67 ENSO-based EWS exist for diseases like malaria, dengue or cholera (Pascual et al 2000, Dhiman et al  
68 2017, Lowe et al 2011; Petrova et al., 2021; Siraj et al., 2014), there is very limited effort dedicated to  
69 zoonotic diseases, particularly in regions beyond the tropics. Importantly, when pinpointing the climate  
70 drivers behind disease transmission, merely establishing associations does not generally fit the purpose  
71 in the current highly changing climate and environmental scenarios, and it therefore becomes crucial to  
72 assemble a mechanistic explanation accounting for the observed effects in terms of epidemic outbreaks.  
73 This is especially true when dealing with delayed, nonlinear synoptic drivers whose cascading  
74 consequences can be intricate and under new conditions not covered by former data. Development of  
75 dynamical models incorporating all these important covariates becomes an optimal way to proceed, in  
76 the context of our current changing climate (Rodó et al., 2013).

#### 77 78 79 80 The North-Atlantic Oscillation

81 In temperate areas, where atmospheric-ocean couplings are weaker, apparent unpredictability has  
82 hindered the use of large-scale weather indices as disease predictors. In particular, the North Atlantic  
83 Oscillation (NAO), a dipole pattern summarized by the surface sea-level pressure difference between  
84 the Subtropical (Azores) High and the Subpolar Low (Barnston and Livezey, 1987; Luterbacher et al.,  
85 2002), is often regarded as containing no climatic memory, in terms of lead time for action. This is  
86 because it has an apparent featureless structure and displays very short memory (at most days or a few  
87 weeks) (Fernandez et al, 2003). The NAO dominates the winter climate variability over the North  
88 Atlantic and its surrounding continents (Hurrell and van Loon 1997). Despite this fact, its behavior in the  
89 low-frequency range and the dynamics behind it are not fully understood yet (Greatbatch 2000; Wanner  
90 et al. 2001; Marshall et al. 2001). Studies based on observational data have shown that extratropical sea  
91 surface temperature (SST) anomalies could force the atmospheric circulation (Czaja and Frankignoul  
92 2002). However, in general, the ocean-atmosphere coupling can lead to a limited enhanced  
93 atmospheric low-frequency variability over the North Atlantic (for reviews see Kushnir et al. 2002;  
94 Visbeck et al. 2003; Czaja et al. 2003). An atmospheric response to midlatitude SST anomalies in  
95 combination with the oceanic gyre circulation are conceived to enhance the decadal variability of the  
96 NAO (Sutton and Allen 1997; Visbeck et al. 1998; Grötzner et al. 1998; Czaja et al. 2003; Paeth et al.  
97 2003). In summary, this variability shows signatures at interannual timescales, as the ones explored in  
98 this study.

#### 99 100 Cutaneous Leishmaniasis in Northern Africa

101 Cutaneous leishmaniasis (CL) is one of the main parasitic vector-borne diseases affecting human beings  
102 and has an estimated annual incidence of around 1 million cases, mainly in the Americas, the  
103 Mediterranean basin, the Middle East and central Asia. CL infections in humans are generally not life-  
104 threatening and often resolve spontaneously, but they give rise to skin lesions and ulcers that can  
105 persist for months or even years, causing significant physical, psychological and social distress among  
106 affected individuals. Additionally, the disease can leave lifelong scars, leading to disability and stigma  
107 (WHO).

108  
109 CL is an emerging disease in Northern Africa. CL was historically limited to several oases located in arid  
110 pre-Saharan zones, where it occurred sporadically leading occasionally to epidemics. However, since the  
111 1980s, the geographical area of CL distribution has spread beyond its desert natural niche and an  
112 increasing number of cases are observed in new endemic foci. Currently, the disease is endemo-

113 epidemic with thousands of cases reported every year in Libya, Algeria, Tunisia and Morocco. (Aoun et  
114 al, 2014)

115  
116 In Northern Africa, CL can be caused by three *Leishmania* species: *Leishmania (L.) infantum*, *L. major*,  
117 and *L. tropica* (Aoun et al 2014). Among these, *L. major* accounts for the highest disease burden and is  
118 responsible for over 90% of the cases reported in Algeria, Tunisia, and Libya. *L. major* is a zoonotic  
119 parasite and is naturally sustained in northern Africa by a two-host cycle: rodents, such as the fat sand  
120 rat (*Psammomys obesus*) or the gerbil (*Meriones* spp).  
121 (<http://www.cfsph.iastate.edu/Factsheets/pdfs/leishmaniasis.pdf>) act as the main reservoirs, and  
122 *Phlebotomus* sandflies (especially *Phlebotomus papatasi*) act as vectors. Different mammals, including  
123 humans, can be infected, but most will present localized lesions and low parasitemia, thereby acting as  
124 dead-end accidental hosts and being unable to further transmit the disease. Hence, CL caused by *L.*  
125 *major* is referred to as Zoonotic CL (ZCL). *L. tropica*, the second most significant *Leishmania* parasite in  
126 North Africa, is on the contrary mostly sustained by anthroponotic transmission chains, thus generally  
127 transmitted between humans and sandflies without involving animal reservoirs (Saik et al 2022). As  
128 compared to Tunisia, Algeria and Libya, the epidemiological situation of CL in Morocco is characterized  
129 by a significantly higher number of cases due to *L. tropica*. This species generates an important disease  
130 burden in the area between the Atlas Mountains and the Atlantic Ocean. The incidence of *L. tropica* is  
131 limited to a few dozen cases in Tunisia, as in Algeria, and the parasite is suspected to be zoonotic with  
132 *Ctenodactylus gondii*, a rocky mountain rodent, acting as the reservoir host (Bousslimi, 2012). Rodent  
133 density, average temperature, cumulative rainfall and average relative humidity have all been shown to  
134 play a role in sustaining and increasing zoonotic CL incidence (Talmoudi et al, 2017). Furthermore,  
135 population acquired immunity may also play a role in modulating CL incidence (Chamakh-Ayari et al,  
136 2017).

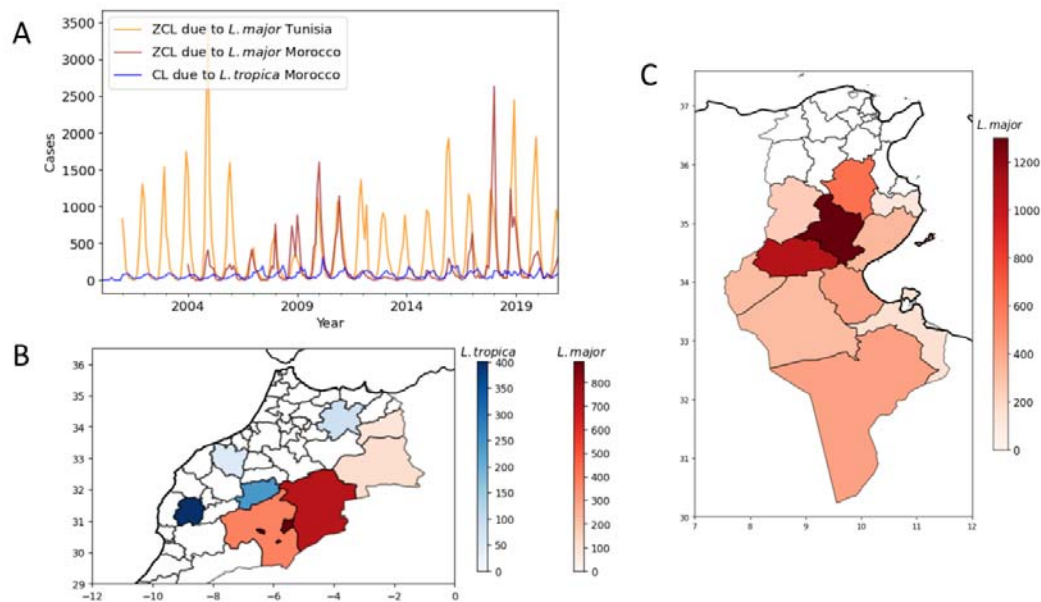
137  
138 This study represents a significant advancement in the attempt to incorporate climate drivers typical of  
139 the northern hemisphere into epidemic-alert systems. It has the goal of deriving a multiscale EWS, by  
140 uncovering the potential for incorporating long-term predictability -up to four years in advance- for  
141 anticipating CL outbreaks in Northern Africa. While keeping with short-term signatures (i.e., the NAO  
142 link to seasonal rainfall), we here added a long-term modulation on the disease epidemiology defined by  
143 the Atlantic Multidecadal Oscillation (AMO), as we uncovered its strong link to CL dynamics in the  
144 region. To substantiate these findings, we furnished a detailed mechanistic pathway that establishes a  
145 direct link between low-frequency climate forcing and the ensuing disease dynamics. Importantly, we  
146 identified a unified, multiscale mechanism that operates both in the short term and over more extended  
147 periods.

#### 148 149 **Data**

150 The incidence of CL was recorded at governorate or provincial scale over a period of 17 to 23 years,  
151 using databases available at regional and national levels in Tunisia and Morocco respectively. Figure 1A  
152 shows the monthly incidence of CL caused by *L. major* in Tunisia along with the incidence of both *L.*  
153 *major* and *L. tropica* infections in several representative Moroccan regions. Figure 1B illustrates the  
154 mean yearly incidence of *L. major* and *L. tropica* in the selected provinces of Morocco. To ensure that  
155 the majority of CL cases are correctly attributed at *Leishmania* species level, specific provinces were  
156 selected: Errachidia, Figuig, Zagora, Jerada, and Ouarzazate being the main transmission hotspots for *L.*  
157 *major*, while Azilal, Chichaoua, Settat, and Taza being the main foci of *L. tropica*. Figure 1C displays the  
158 mean yearly incidence of CL caused by *L. major* in Tunisia at the governorate level.

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Fig 1: (A) Monthly incidence of *L. major* infections in Tunisia and in studied provinces of Morocco, along with the incidence of *L. tropica* infections in studied regions of Morocco. (B) Mean incidence at the province level of *L. tropica* and *L. major* in Morocco (zoomed to regions with data). (C) Mean incidence at the province level of *L. major* in Tunisia.

167 Reported cases of CL in North Africa show a seasonal distribution (Figure 1A). Notably, *L. major* ZCL  
168 exhibits a consistent seasonal occurrence in both Tunisia and Morocco, with case records peaking  
169 around December. This seasonal pattern of ZCL is closely related to *Phlebotomus papatasi* phenology in  
170 the North African region, with sandfly adult activity starting in May–June and extending to October.  
171 There is a long-lasting peak of sandfly density 3 to 3.5 months before the peak of CL incidence (Chelbi et  
172 al, 2007). The peak of *L. tropica* CL incidence around March is also likely attributed to sandfly vector  
173 phenology but also to a longer *L. tropica* incubation period compared to that of *L. major* leading to  
174 delayed onset of lesions (Aoun et al, 2022). On the other hand, Figure 1A also shows that CL incidence  
175 displays clear inter-annual fluctuations. This is particularly evident with ZCL caused by *L. major*, for  
176 which seasonal peaks varied from hundreds to thousands of cases according to the year (Figure 1A).

177

## 178 Results

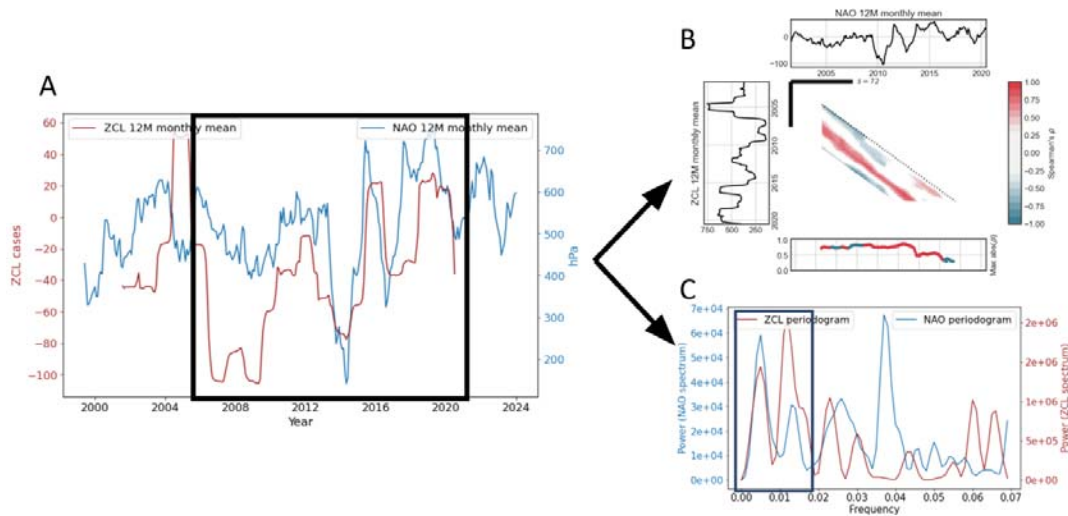
### 179 Multiscale analysis of the NAO

180 Precipitation in northern Africa appears to be essentially modulated in the wet season by the North  
181 Atlantic Oscillation (NAO) and its associated pressure systems. Such pressure differences modulate the  
182 strength of the North-Atlantic westerly winds, which are responsible for the transport of moisture  
183 across the sea, and therefore impact both precipitation and temperature in extensive regions of Europe  
184 and Northern Africa (Toumi et al, 2012). CL is known to be impacted by such synoptic weather patterns,  
185 especially through modulation in precipitation. As such, we assessed whether the NAO index provided  
186 any skill that could be used in accounting in advance for CL dynamics.

187

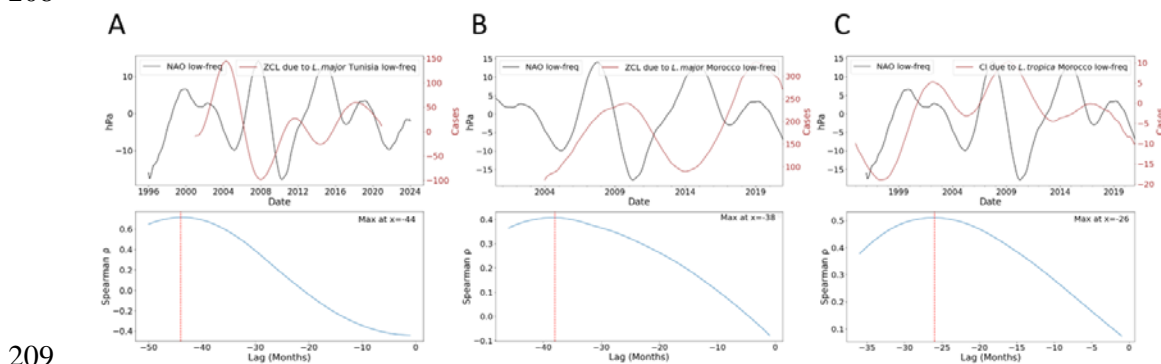
188 Analysis of the NAO effects on ZCL in Tunisia revealed strong and clear lagged low-frequency couplings  
189 existing between NAO and ZCL incidence. Fig 2A displays both the NAO index and the ZCL incidence in  
190 Tunisia, both smoothed with a rolling mean to remove the short-term variability. When lagged around -  
191 44 months (3.5 years), a very robust coupling emerges, which is maintained throughout the whole  
192 period. Fig 2B shows how stable the coupling is, by performing local-lagged Spearman correlations  
193 between the time series in small windows by running a Scale-Dependent Correlation (SDC) Analysis. Fig

194 2C shows on the spectral domain this same match between dominant periods at the low-frequency end  
 195 of the spectrum, thereby reinforcing the aforementioned findings.  
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198  
 199 Fig 2: (A) ZCL incidence in Tunisia and NAO index, smoothed with a 12-M moving average and overlaid after lagging  
 200 them -46 months. (B) SDC-Analysis of the two time series showing consistent strong couplings throughout the  
 201 whole period. (C) Periodogram of both time series showing the match at the low-frequency end of the spectrum.  
 202

203 To further study the low-frequency coupling existing between the NAO index and CL, a Singular  
 204 Spectrum Analysis (SSA) was applied to both the NAO and the three country-wise aggregated CL time  
 205 series (*L. major* in Tunisia, *L. major* in Morocco and *L. tropica* in Morocco). Of relevance, a common low-  
 206 frequency component emerged from the four aforementioned SSA decompositions with a period of  
 207 around 7-10 years (Figure 3).  
 208



209  
 210 Fig 3: Low-frequency coupling NAO-CL. (A) For ZCL due to *L. major* in Tunisia. Up: Overlaid low-frequency  
 211 components. Down: Cross-correlation analysis, showing a peak at -44 months ( $p$ -value<0.05). (B) For ZCL due to *L.*  
 212 *major* in Morocco. Up: Overlaid low-frequency components. Down: Cross-correlation analysis, showing a peak at -  
 213 38 months ( $p$ -value<0.1;  $n$  being shorter) (C) For CL due to *L. tropica* in Morocco. Up: Overlaid low-frequency  
 214 components. Down: Cross-correlation analysis, showing a peak at -26 months ( $p$ -value<0.05).  
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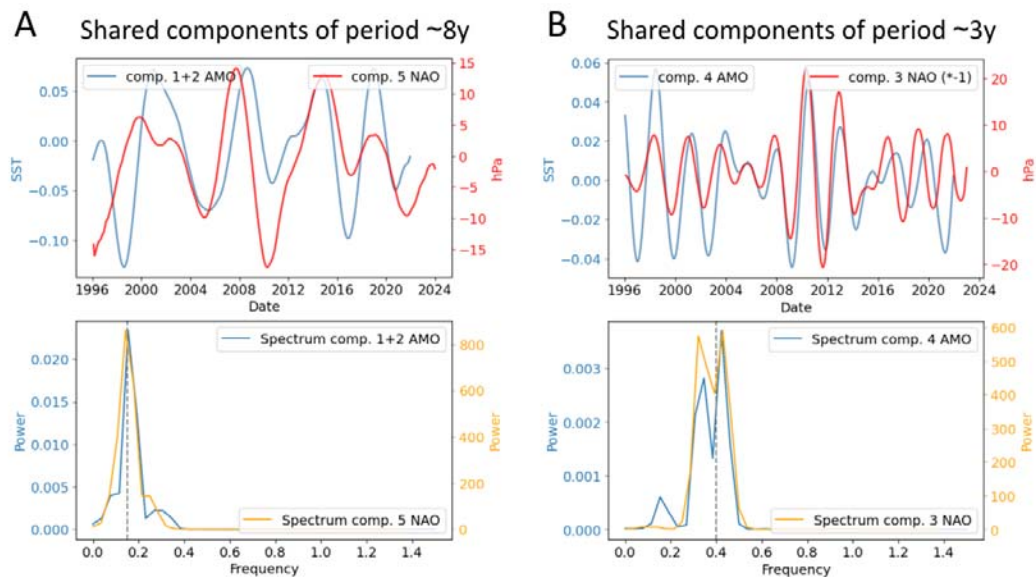
216 For the three CL situations, cross-correlation analyses between the low-frequency NAO and the low-  
 217 frequency CL time series resulted in a unimodal function with a peak at -44 months (Tunisia *L. major*,  $p$ -  
 218 value<0.05), -38 months (Morocco *L. major*,  $p$ -value<0.1;  $n$  being shorter) and -26 months (Morocco *L.*  
 219 *tropica*,  $p$ -value<0.05) (Fig. 3). Similar lags, of around 3-3.5 years therefore naturally arose in Tunisia and  
 220 Morocco for *L. major*. In parallel, a strong NAO-*L. tropica* match also emerged, despite the latter being  
 221 at shorter lags.  
 222

223 Given the long lead times obtained, we proceeded to investigate the role and associated signatures of  
 224 the Atlantic Multi-Decadal Oscillation (AMO), a coherent mode of interdecadal climate variability

225 occurring in the North Atlantic Ocean with a fundamental estimated period of 60-80 years, based upon  
 226 the average anomalies of SSTs in the North Atlantic basin.

227  
 228 To this end, a SSA was again computed but now for the AMO index to have it decomposed into its main  
 229 modes of variability. Cross-inspection with the former SSA executed over the NAO showed common  
 230 low-frequency covariating components between the two phenomena. In fact, such decompositions  
 231 clearly display that the NAO-AMO coupling effectively operates synchronously across various temporal  
 232 scales. The characteristic period of 7-10 years previously found in all CL time series and the NAO index,  
 233 similarly shows up in the AMO (Figure 4A). Additionally, it became clear there exists a strong coherency  
 234 between the 3rd component of the NAO and the 4th component of the AMO index, both showing a  
 235 similar frequency peak at periods between around 2 - 3.5 years (Fig. 4). Collectively, these components  
 236 elucidate approximately 17% and 7% of the AMOs and NAOs original variance, respectively.

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239 Fig 4. NAO-AMO coupling. (A) Shared components with a period of around 8y. (Up) Superposition of the 5th NAO  
 240 component and the 1+2 AMO components (Spearman's  $\rho=0.48$ ,  $p\_value<0.05$ ). (Down) Overlaid Spectral  
 241 decomposition of such NAO and AMO 8y components (B) Shared components of period of around 3y (Up)  
 242 Superposition of the 3rd NAO component with the 4th AMO component (Spearman's  $\rho=0.53$ ,  $p\_value <0.05$ ).  
 243 (Down) Overlaid spectral decomposition of such NAO and AMO 3y components.

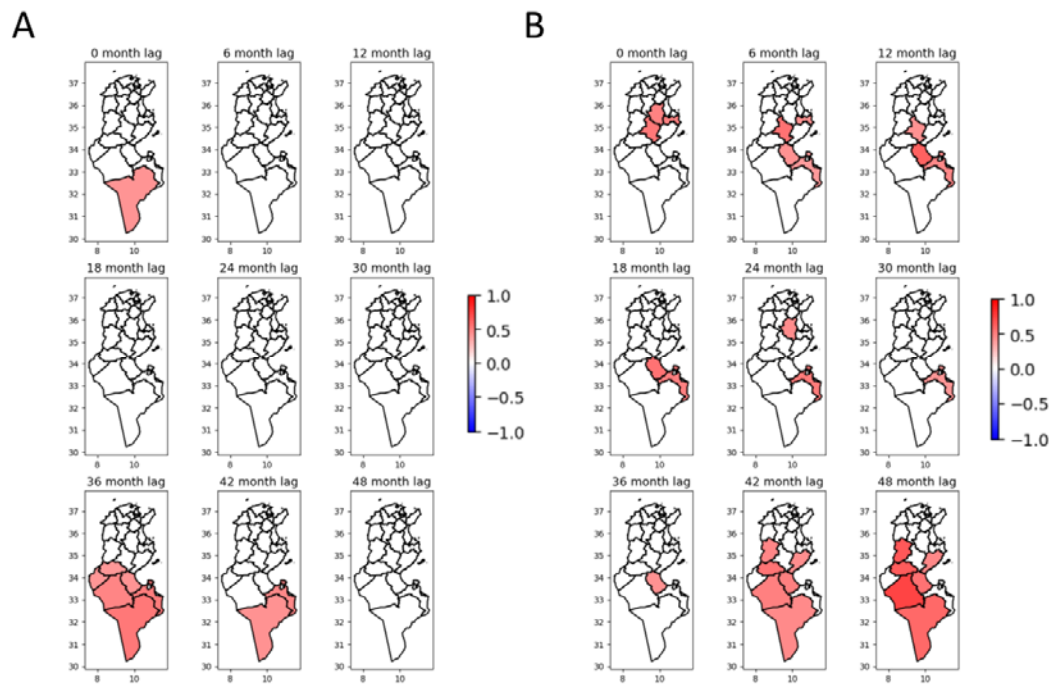
244  
 245 The overlaid ~8-year reconstructed components of both the leishmania parasites and the climate indices  
 246 under investigation (namely, the NAO, AMO, *L. major* Tunis, *L. major* Morocco, *L. tropica* Morocco) can  
 247 be seen in Fig S1.

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### Short-term responses

252 As previously discussed, the NAO modulates the strength of the westerly winds, which are responsible  
 253 for carrying moisture across the Atlantic, and therefore they seasonally impact both precipitation and  
 254 temperature in Europe and Northern Africa. It is furthermore surprising to find a large lagged effect  
 255 given that the NAO instead has an instant effect on precipitation and quickly loses predictability  
 256 (Wunsch, 1999). Indeed, at these very short time scales ZCL in Tunisia is seen to be strongly associated  
 257 with synoptic weather events displaying a dipole linked to the NAO phenomenon through the Açores  
 258 pole (Fig S3), despite being also partly related to the Scandinavian Mode (Wang and Tang, 2020). Its  
 259 associated skill can be traced back to a maximum of 6 months in the form of leading sea-level pressure  
 260 anomalies (Fig S3).

261 To further delve into the climatic forcing mechanisms operating on ZCL dynamics, we zoomed in and  
262 regionalized the analyses in Tunisia. We identified an expected significant short-term coupling between  
263 the NAO and precipitation raw values (not shown), which also showed up after applying a 12-month  
264 smoothing filter (Fig 5A). However, apart from the short-term coupling, the associations at longer lags  
265 previously identified similarly showed up (Fig 5A).  
266  
267



268 Fig. 5: (A) Significant full series correlations ( $p$ -value  $< 0.05$ ) of NAO vs. Rainfall (smoothed with 12-mo moving  
269 windows) in Tunisia at the province level from 2007 to 2021. (B) Significant correlations ( $p$ -value  $< 0.05$ ) NAO vs ZCL  
270 (smoothed with a 12-mo moving average) in Tunisia at the provincial level for the period 2007-2021  
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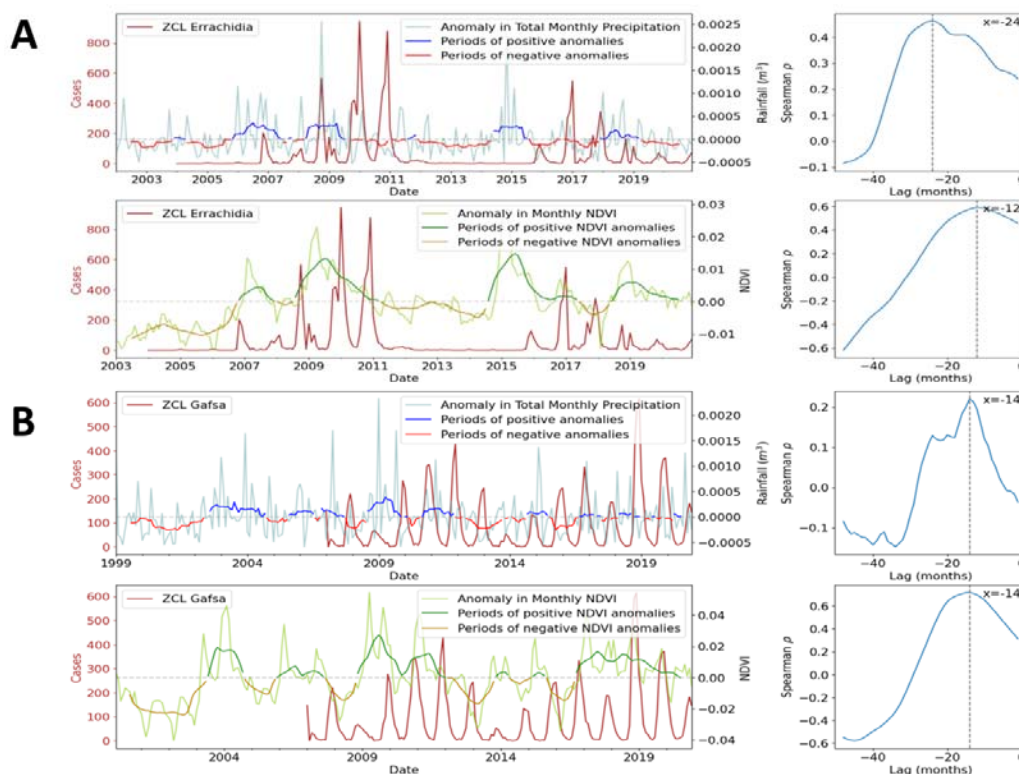
274 In the previous analysis a maximum low-frequency NAO-ZCL coupling was found for Tunisia with a 44  
275 months delay (Fig 2A). We now obtained a maximum NAO-Precipitation coupling located at 36 months  
276 delay, therefore around 1 year before that obtained for the NAO-ZCL, thereby reinforcing our previous  
277 findings and showcasing how precipitation mediates in the cascade effects of NAO on ZCL in Tunisia.  
278

279 We proceeded to perform the same analyses but, now instead, directly between the NAO and ZCL. In a  
280 higher spatial resolution domain, we again obtained evidence regarding the existence of a strong NAO-  
281 ZCL coupling at around -44 months lag, manifesting especially in the middle and southern part of  
282 Tunisia, where we also saw the best couplings with precipitation. (Fig 5A, 5B)  
283

284 The relationship between NAO and precipitation in Morocco is less direct and evident compared to the  
285 situation in Tunisia (Figs S5, S6).  
286

287 Local precipitation impacts ZCL short-frequency dynamics in both countries (Fig. 6). At the Governorate  
288 level / Wilaya level, ZCL and precipitation tend to show a maximum coupling at around a year lag.  
289 Interestingly, the normalized difference vegetation index (NDVI), known to respond tightly to rainfall in  
290 these types of climates, displayed higher correlations with ZCL at slightly shorter lags. Figure 9 shows a  
291 noticeable example of such coupling for Errachidia (Morocco) and Gafsa (Tunisia).  
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Fig. 6. ZCL-Rainfall couplings (up) and ZCL-NDVI couplings (down) -and cross-correlation analyses (CCF), right- for (A) Errachidia (Morocco) and (B) Gafsa (Tunisia). All original rainfall, NDVI and ZCL time series are shown (in light blue, light green and brown). The graphs also display the result of 12-mo moving filter on the climatic and environmental time series, which serve to illustrate periods of positive anomalies (in blue for precipitation and green for NDVI) and negative anomalies (in red for precipitation and yellow for NDVI). CCFs are performed after applying a 12-mo moving filter to all time series to remove short term variability. The correlation peaks for the cross-correlations from Errachidia all pass a significance test at  $\alpha=0.05$ . For Gafsa, and mainly because the timeseries is shorter, the cross-correlation peak with rainfall doesn't pass a test ( $p_{\text{value}} < 0.1$ ) but it does with NDVI.

## Discussion

Very few studies have attempted to investigate the inter-annual modulation of CL and associate it to external forcings. One seminal study identified 3y cycles in CL data from Costa Rica and hypothesized them to be connected to the El Niño Southern Oscillation (ENSO) phenomenon (Chaves et al, 2006). A shared frequency peak was identified between CL incidence and the ENSO MEI index, but the study did not provide a detailed mechanistic explanation of how the atmosphere-ocean coupling takes place from the source region in the equatorial SST and the local CL dynamics in Costa Rica. As such, the exact mechanistic cascading pathway underlying the findings remained elusive.

In this study, instead, we proved that despite the distance between Morocco and Tunisia, and the large known differences between *L. major* and *L. tropica* epidemiological cycles, all the studied different CL disease situations in North-Africa show strong low-frequency components being shared with the NAO. In fact, such couplings take place at very similar lags for the two *L. major* situations, but at a different lag for *L. tropica*. This apparent discrepancy should not be surprising and reinforces these findings. *L. tropica* represents a non-zoonotic form of CL, that does not appear to operate through an animal reservoir. Therefore, we hypothesize that it presumably responds faster to the same extrinsic climatic stressor, which would explain the shorter lags found. Overall, we inferred that CL inter-annual dynamics are governed by a large-scale climate synoptic pattern, which likely operates along with population acquired immunity to determine the final low-frequency patterns of the CL time series (Chamakh-Ayari et al, 2017).

Overall, we identified two shared components between the NAO and the AMO illustrating that both are showing responses to the same climatic process. Therefore, from a mechanistic perspective, a unique

329 unified mechanism can explain both the multiscale short-term and long-term relationships between  
330 climate and ZCL epidemiology in Northern Africa.

331  
332 Regarding the exact pathways by which such large-scale indices locally affect transmission, we identified  
333 a strong modulation of precipitation in Tunisia, both instantaneous and at 3 years lag. On the latter it is  
334 important to say that such a 3-year lag is around one year before the maximum couplings appear with  
335 respect to the disease (which manifest around a 4-year lag). This one-year difference between the  
336 cascading effects of precipitation on the disease (i.e., before effects show up in ZCL incidence) was also  
337 independently identified in Fig 9, thereby reinforcing our findings. Furthermore, by integrating data  
338 from vegetation, which strongly couples to ZCL incidence, we found evidence that supports the notion  
339 that vegetation plays a key role in sustaining and increasing zoonotic CL incidence. These results are in  
340 line with those of other studies showing that the incidence of ZCL in central Tunisia increased  
341 significantly when there was an increase in rainfall lagged by 12 to 14 months (Toumi et al, 2012). They  
342 are also in accordance with results of other studies that found that the occurrence of ZCL is related to  
343 rodents' density but lagged 2 months (Talmoudi et al, 2017). In fact, higher rainfall is expected to result  
344 in increased density of vegetation, particularly for chenopods, a halophytic plant that constitutes the  
345 exclusive food source of *P. obesus*. Presumably, following a high density of *P. obesus*, the source of *L.*  
346 *major* transmissible from the rodents to blood-feeding female sand flies could lead to a higher  
347 probability of transmission to humans over the next season.

348  
349 In Morocco, the NAO did not so clearly appear to modulate precipitation, especially when aggregating  
350 data from the whole country. Morocco's unique and complex geographical position, being in a fringe  
351 topographic region separated by the Atlas Mountains, may be behind the fainter signature -albeit also  
352 significant- obtained with CL. This mountain range is responsible for the distinct manifestation of  
353 precipitation patterns on either side, ultimately being influenced by diverging interactions with the  
354 Atlantic moist winds and steep topography. The nonlinearities in the observed Morocco's ZCL response  
355 to the NAO could also be due to these more complex, multifaceted pathways by which the NAO  
356 modulates climate in northern Africa.

## 357 358 **Conclusions**

359  
360 The AMO low-frequency oceanic pattern in the North Atlantic modulates the low-frequency dynamics of  
361 ZCL epidemiology in North Africa. Its fundamental role can explicitly be traced to the cascading effects  
362 connecting the long-term SST changes in the North Atlantic to the ZCL dynamics. In particular, we could  
363 show how some of the main modes of variability of the AMO are transferred via the NAO. Physically,  
364 this means that variations in the SST anomalies of the Atlantic Ocean may operate in the form of large-  
365 scale changes in synoptic weather conditions subsequently affecting the moist westerly winds patterns  
366 over the Atlantic. This in turn creates changes in precipitation across North Africa that can be traced to  
367 ultimately affect CL incidence through cascading effects in vegetation.

368 A multi-scale prediction system such as the one described in this study, that can operate at both long  
369 (roughly 4 years) and short (6 months) having the NAO and the AMO as the sources of climatic memory,  
370 therefore predictability, has never been built before. Importantly our results identify a unique multiscale  
371 mechanism accounting for both short-term and long-term dynamics, that could be promptly integrated  
372 to provide an alert system for the different types of cutaneous leishmaniasis in the region.

## 373 374 375 **Methods**

376  
377 Scale-Dependent Correlation (SDC) analysis is an optimal method for identifying dynamical couplings in  
378 short and noisy time series (Rodo et al, 2006). In general, Spearman correlations between incidence and  
379 a meteorological time series assess whether there is a monotonic relation between the variables. SDC  
380 analysis was specifically developed to study transitory associations that are local in time at a specified  
381 temporal scale corresponding to the size of the time intervals considered ( $s$ ). The two-way  
382 implementation (TW-SDC) is a bivariate method that computes non-parametric Spearman rank  
383 correlations between two time series, for different pairs of time intervals along these series. Different  
384 window sizes ( $s$ ) can be used to examine increasingly finer temporal resolution. Significance is assessed  
385 with a non-parametric randomization test. For the baseline test, SDC calculates Spearman correlations

386 (at  $\alpha=0.05$ ) between two white-noise time series at each fragment size  $s$  for a non-parametric  
387 permutation test. The indices of the series are randomly re-ordered, breaking their temporal shape. This  
388 permutation test enables a first estimate of the probability of finding significant spurious correlations,  
389 and it can thus be used as a non-parametric significance test for pairs of any length for the time series of  
390 interest. (The code in Python for Scale-Dependent Correlation Analysis (SDC) developed by X.R. can be  
391 found at <https://github.com/AlFontal/sdcpy>.)  
392

393 Singular spectrum analysis (SSA) (Ghil et al 2002, Vautard et al 1992) was applied to separate different  
394 orthogonal components in both the CL time series and the climate ones. SSA involves the spectral  
395 decomposition (eigenvalues and corresponding eigenvectors) of a covariance matrix obtained by lagging  
396 the time series data for a prescribed number of lags  $M$  called the embedding dimension. There are two  
397 crucial steps in this analysis for which there are no formal results but useful rules of thumb: one is the  
398 choice of  $M$ ; the other is the grouping of the eigenvectors to define the specific major components and  
399 reconstruct them. Typically, the grouping of the eigencomponents is based on the similarity and  
400 magnitudes of the eigenvalues, their power (variance of the data they account for), and the peak  
401 frequency of resulting reconstructed components (RC). For the selection of the embedding dimension  
402 one general strategy is to choose it so that at least one period of the lowest frequency component of  
403 interest can be identified, that is  $M \geq 2\pi f_s / f_r$ , where  $f_s$  is the sampling rate and  $f_r$  is the minimum  
404 frequency. Another strategy is that  $M$  be large enough so that the  $M$ -lagged vector incorporates the  
405 temporal scale of the time series that is of interest. The larger the  $M$ , the more detailed the resulting  
406 decomposition of the signal. In particular, the most detailed decomposition is achieved when the  
407 embedding dimension is approximately equal to half of the total signal length. A compromise must be  
408 reached, however, as a large  $M$  implies increased computation, and too large a value may produce a  
409 mixing of components. SSA, given its data-adaptive nature and orthogonal decomposition approach, can  
410 be used as an alternative smoothing method to moving averages, which is more flexible and robust, and  
411 also deals better with non-stationary noisy time series (Ghil et al., 2002).

412 Rainfall in semi-arid regions typically exhibits a lot of stochasticity. We applied moving averages in some  
413 situations to extract lower-frequency components. In such instances, and also in the analysis of low-  
414 frequency components, given that both series become more autocorrelated and their number of  
415 degrees of freedom decreases, we calculated  $p$ -values by correcting by the effective degrees of  
416 freedom. In particular, we employed a bootstrap approach to assess their significance. To compute the  
417 distribution of resulting correlation coefficients from which we determined significance, we repeatedly  
418 randomized the smoothed time series while keeping original autocorrelation.

419 To study responses in vegetation to rainfall variability ultimately impacting CL, we used the normalized  
420 difference vegetation index (NDVI). NDVI is a remote sensing measurement used to assess the density  
421 and health of vegetation. It calculates the difference between near-infrared (which vegetation strongly  
422 reflects) and red light (which vegetation absorbs) from satellite or aerial imagery, providing a numerical  
423 indicator that can be used to monitor plant growth, vegetation cover, and biomass production.

424  
425

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